

Beyond Equilibrium

Reaction Kinetics in Geodynamic Models

Collab:

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Evidence of disequilibrium in rocks



Meta-igneous (anorthosite), Bergen Arcs, Holsnøy Norway

plagioclase + pyroxene → garnet

Solid-state reactions in rocks

Thermodynamics

Q: What *should* happen?

Kinetics

Q: Does it *actually* happen on geodynamic timescales?

Aspirational goal for ASPECT

Go beyond equilibrium:

Parameterize reaction progress in a physically meaningful and computationally tractable way

The basis of kinetic theory

Transformation rate of the reaction:



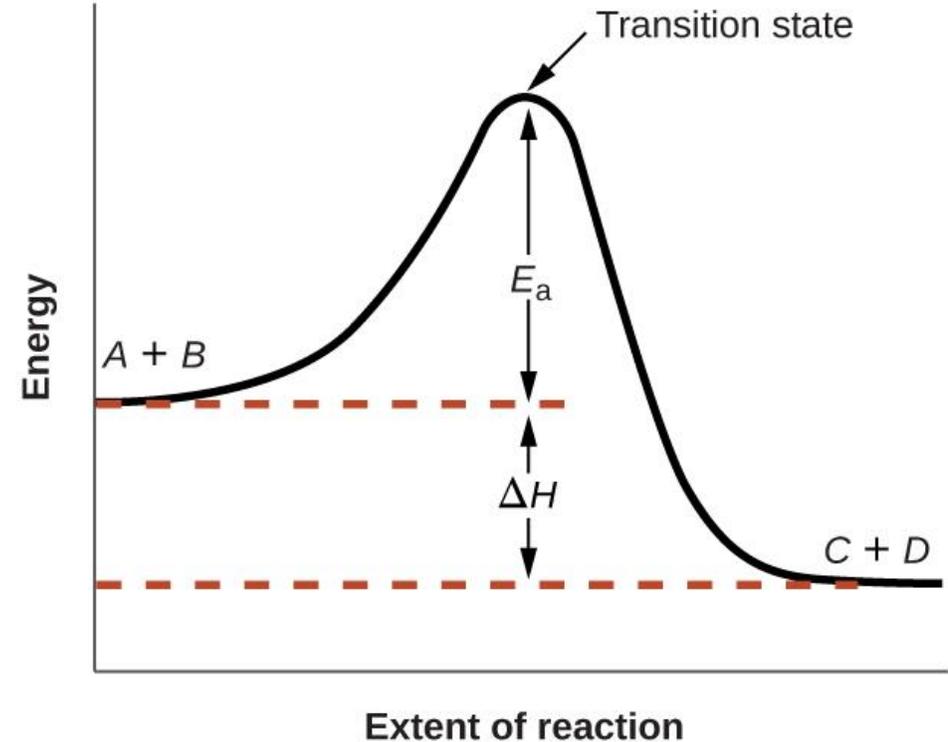
depends on the number of atoms that can achieve the transition state

$$\frac{dy}{dt} \propto A \exp\left(\frac{-E_a}{RT}\right)$$

y : volume fraction transformed

A : frequency factor

E_a : activation energy



The basis of kinetic theory

A more general form of transformation rate is:

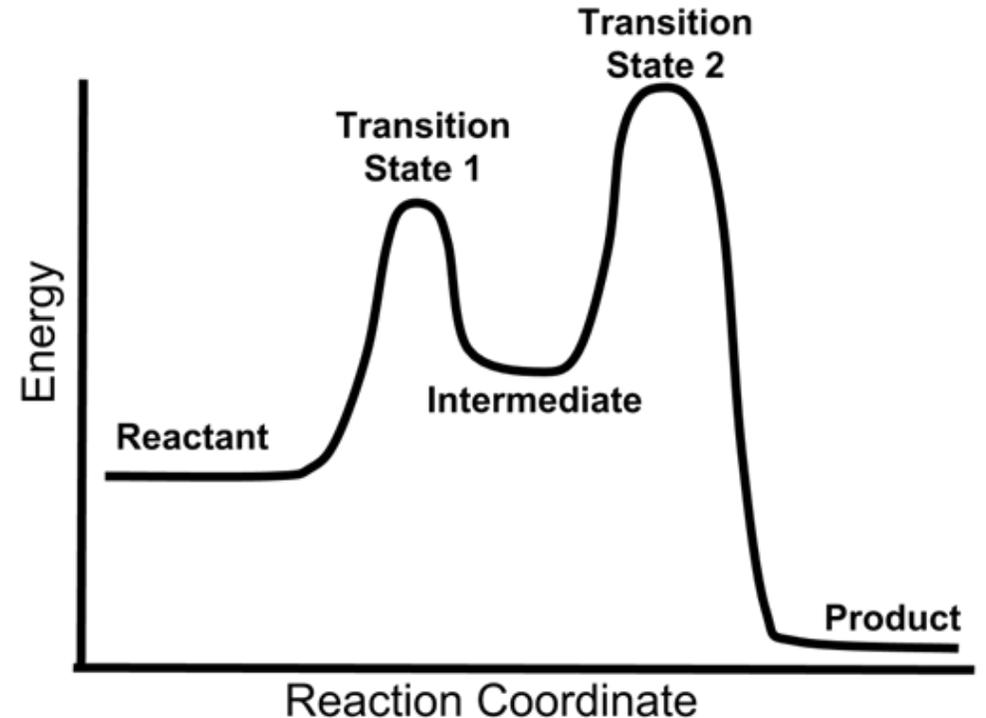
$$\frac{dy}{dt} = k(P, T) f(y)$$

k : rate constant (**pathway dependent**)

$$k = A \exp\left(\frac{-E_a}{RT}\right)$$

$f(y)$: how reaction progress feeds back on rate

- availability of reactants
- evolving interface area
- site saturation
- transport distances



Mechanisms and pathways

Reaction mechanism: how atoms move

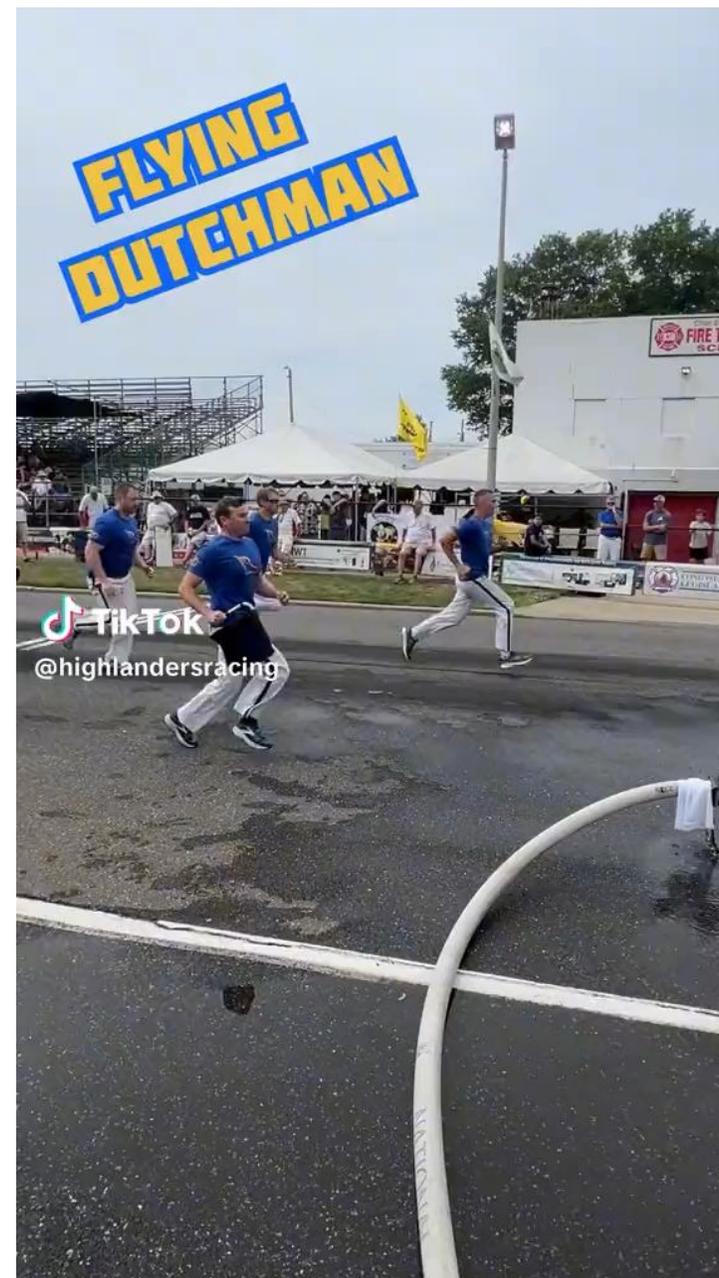
Reaction pathway : sequence of steps (mechanisms) to complete reaction

- Can be complex
- Slowest step limits overall reaction rate*

*If pathway is sequential

Example:

dissolution → grain-boundary diffusion → precipitation
scoop bucket → move bucket up ladder → pour bucket



JMAK model

Johnson–Mehl–Avrami–Kolmogorov (1937–1941)

Many reactions $\alpha \rightarrow \beta$ generally follow:

$$y = 1 - \exp(-kt^n)$$

k : rate constant (Arrhenius)

n : Avrami exponent*

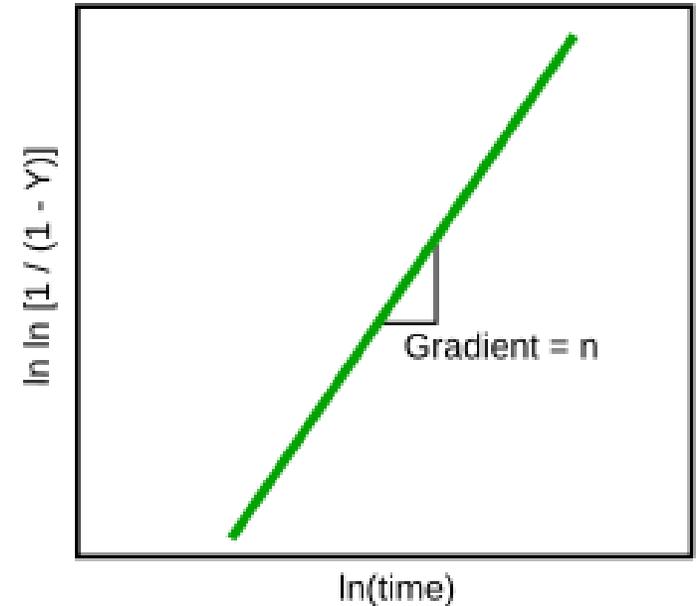
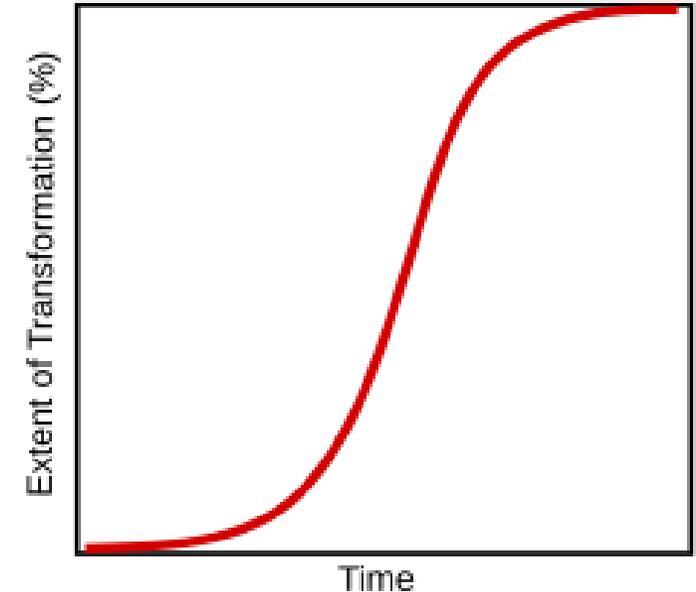
*Interpreted as relating to geometry + mechanism

$n = 4$: nucleation + 3d growth (spheres)

$n = 3$: 3d growth (spheres)

$n = 2$: 2d growth (plates)

$n = 1$: 1d growth (needles)



JMAK model

Updated for nucleation at preferred sites (Cahn, 1956):

$$y = 1 - \exp\left(-\frac{1}{3}\pi\dot{N}\dot{G}^3t^4\right)$$

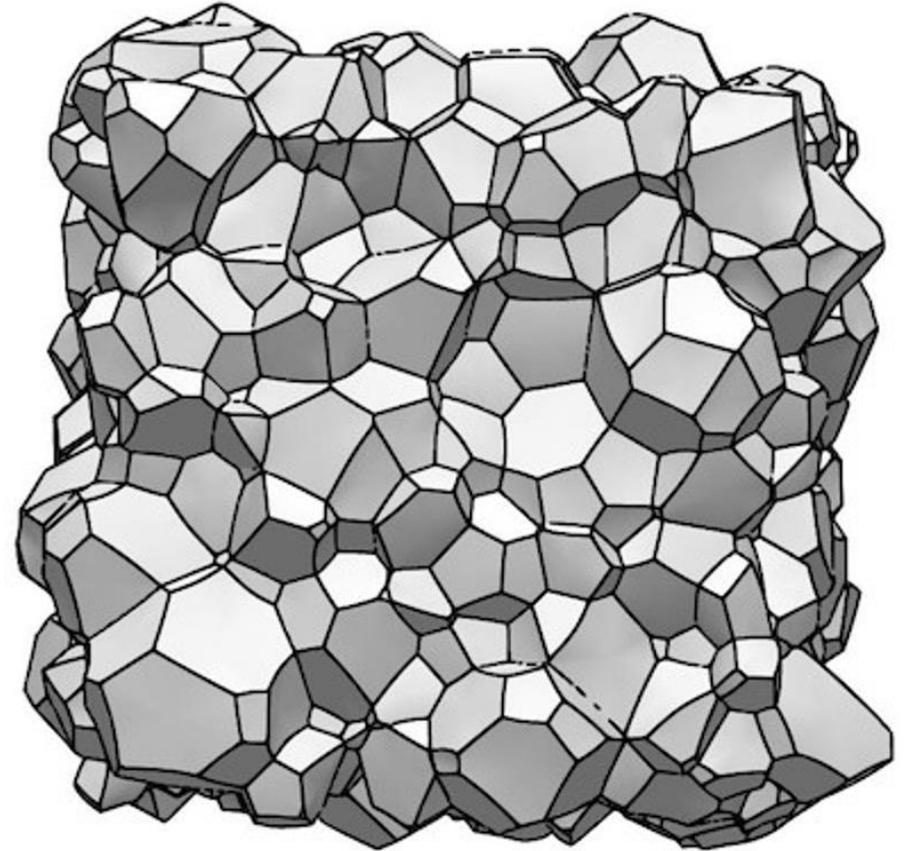
\dot{N} : nucleation rate per volume ($\text{m}^{-3} \text{s}^{-1}$)

\dot{G} : growth rate of grain boundary (m s^{-1})

After nuclei saturate grain surfaces:

$$y = 1 - \exp\left(-\frac{6.7}{d}\dot{G}t\right)$$

d : mean grain size



JMAK model

After nuclei saturate grain surfaces (Cahn, 1956):

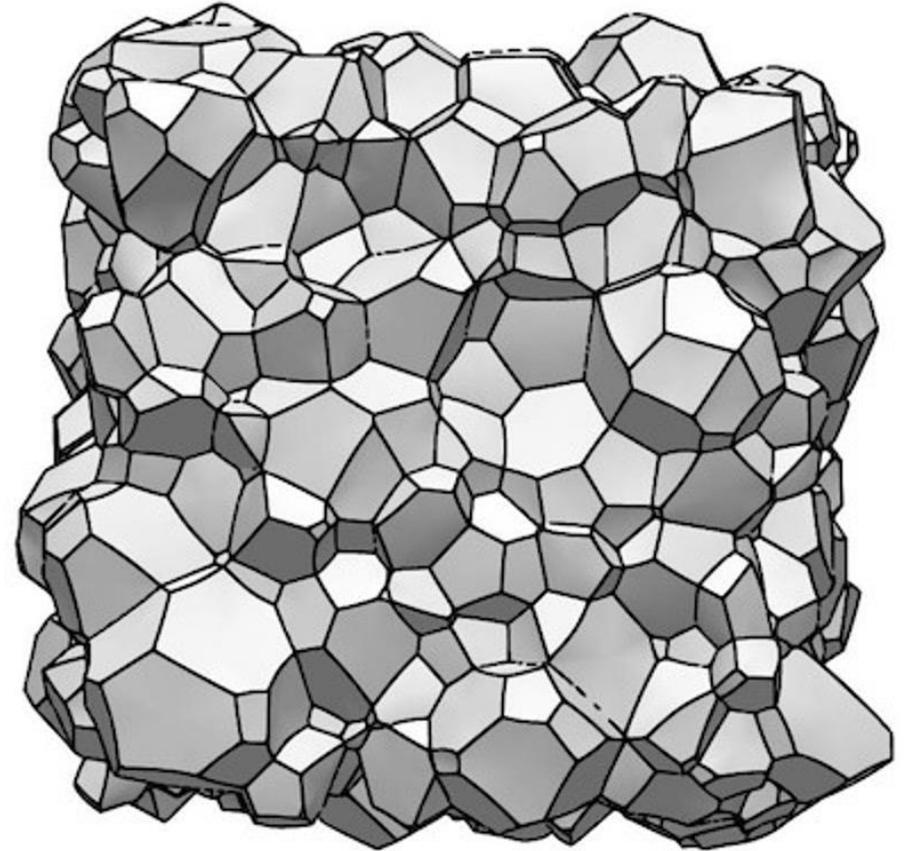
$$y = 1 - \exp\left(-\frac{6.7}{d} \dot{G} t\right)$$

If \dot{G} is constant at fixed PT (no long-range diffusion):

$$\frac{dy}{dt} = \frac{6.7}{d} \dot{G} (1 - y)$$

where \dot{G} follows basis of kinetic theory:

$$\dot{G} = A T \exp\left(\frac{-E_a}{RT}\right) \left[1 - \exp\left(\frac{-\Delta G_{rxn}}{RT}\right)\right]$$



Diffusion-controlled growth

If grain growth is not constant at fixed PT (long-range diffusion):

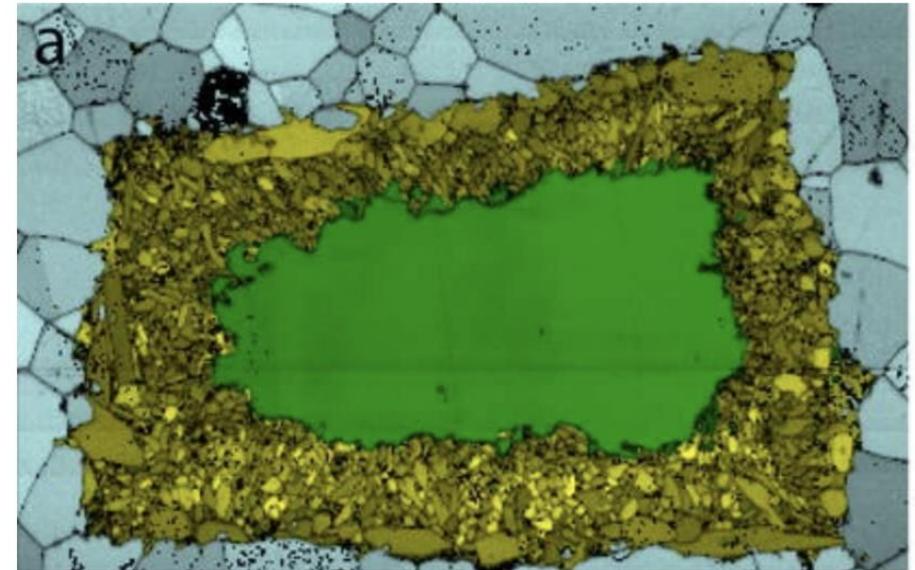
$$x \propto \sqrt{Dt}$$

$$\dot{G} \propto \frac{1}{\sqrt{t}}$$

then we have a different numerical problem altogether*

$$y = 1 - \exp(-kt^{1/2})$$

*Still tractable without explicitly tracking rim width x (e.g., Dähler & Yuen 1996; Reidel & Karato 1997)



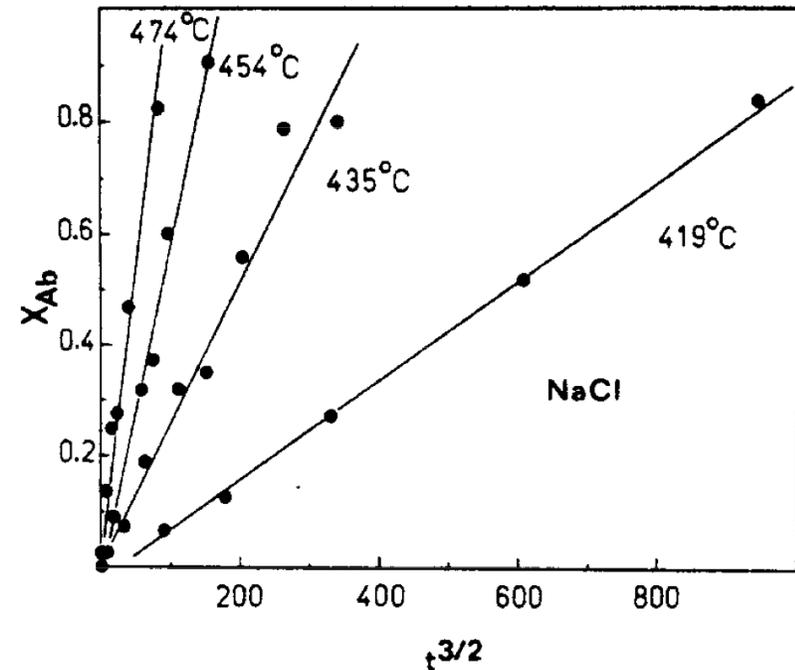
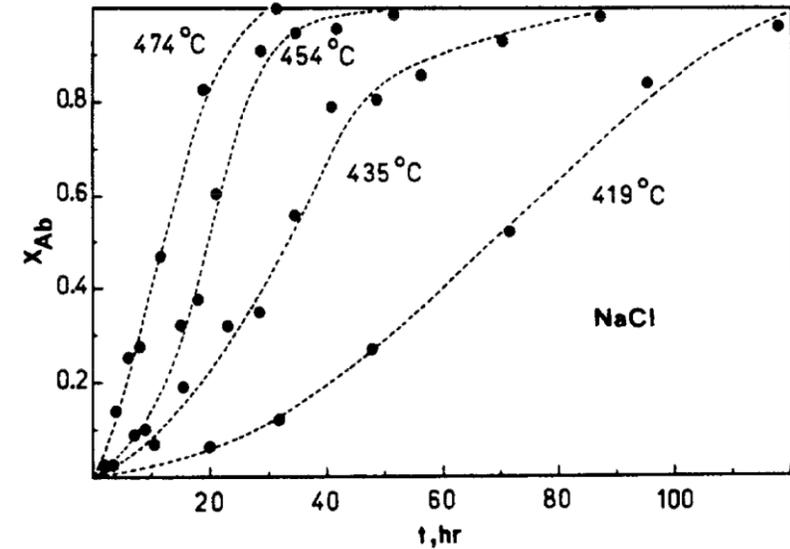
Simpler models

Power law:

$$y = kt^n$$

ex: zeolite + quartz \rightarrow albite + H₂O

Coefficients determined by measuring y during isothermal experiments at different T



Summary

The following rate laws could be considered for implementation in ASPECT***

Constant rate:	$y = kt$	simple; limited applications
Power law:	$y = kt^n$	simple; no theoretical basis; limited data
JMAK:	$y = 1 - \exp(-kt^n)$	nucleation + growth; many applications; limited data
JMAK (Cahn):	$y = 1 - \exp\left(-\frac{1}{3}\pi\dot{N}\dot{G}^3t^4\right)$	nucleation + growth; many applications; limited data
JMAK (Cahn):	$y = 1 - \exp\left(-\frac{6.7}{d}\dot{G}t\right)$	growth; many applications; limited data

*****Requires many assumptions** about reaction pathways, microstructures, grain sizes, and more ...

Ex: Li et al. (2026 in prep) Kerswell et al. (2026 in review)